This work presents a case study for the lakes Barun-Torey and Zun-Torey, which have been subject to major changes in water level. Both lakes are located in the territory of the Daursky State Nature Biosphere Reserve (acrid steppe zone of the Palaeartic region, South-East Transbaikalia, Russia), in the border area with Mongolia. The Torey lakes are large and shallow. A unique feature of the Torey lakes is their unstable hydrological regime. The water level variation is explained by the periodic filling and drying of the lakes due to cyclical climatic changes in humidity and temperature. A brief overview of some hydrological and hydrochemical characteristics of the lakes, using references and unpublished (own) data, is provided. The lakes have been studied for a number of years, which have been provisionally divided into some periods. There are three periods for Lake Zun-Torey: high (1999 and 2003), intermediate (2007, 2011) and low (2014 and 2016) lake levels, and two periods for Barun-Torey: high water level (1999 and 2003) and the initial filling phase (2014, 2016, 2018, 2020). The aim of this paper is to synthesise data on changes in the environmental parameters of the Torey lakes during a climate cycle (the drying and initial filling of basins). Climatic fluctuations lead to changes on some physical parameters and the hydrochemical compound. In Lake Zun-Torey, such environmental characteristics are total dissolved solids, pH, and water temperature increase, while depth decreases during transitions from high to low water levels. In periods of high water level, the hydrochemical parameters of Lake Barun-Torey are the same as those of Lake Zun-Torey (oligohaline and alkaline). The temporary pools on the dry bed of Lake Barun-Torey are freshwater, arising from groundwater and precipitation.

**Key words:** abiotic parameters, Daursky State Nature Biosphere Reserve, hydrological cycle, Lake Barun-Torey, Lake Zun-Torey, water level

**Introduction**

Salt lakes are geographically widespread, numerous and a significant part of the world’s inland aquatic ecosystems (Hammer, 1986). They are important natural assets with considerable aesthetic, cultural, economic, recreational, scientific, conservation and ecological values (Zadereova, 2013). Saline lakes develop as the termini of inland drainage basins where hydrological inputs and outputs are balanced. These conditions occur in arid and semi-arid regions (Williams, 2002). Shallow soda lakes attract considerable attention because of their variable and individualistic nature (Egorov, 1993; Tóth et al., 2014; Schagerl, 2016; Boros & Kolpakova, 2018), and are sensitive to environmental changes, especially changes in the precipitation and strongly dependent on the hydrological budget. Alternation of dry and wet periods causes changes in the physical (water temperature, water level), and chemical (extensive seasonal and inter-annual variations in salinity and pH) characteristics of such a lake (Comín et al., 1992; García et al., 1997; Shadrin & Anufriieva, 2013).

Dauria, located in the northern part of Central (Inner) Asia, is counted among those regions which ecological state is particularly dependent on climatic change. Most of the Daurian ecoregion lies in the north and northeast of China and Eastern Mongolia but it also includes the Russian forest-steppe and steppe regions of the Zabaikalsky Kray and the Republic of Buryatia (Kirilyuk & Tkachuk, 2012). Dauria is characterised by intrasecular cycles, lasting from 8–10 to 35 years. The last complete cycle of changes in atmospheric moisture regime in this territory lasted 34 years (from 1965 to 1998) and had equal periods of aridity (to 1981) and humidity (Zamana & Borzenko, 2010; Bazhenova, 2013). From 1982 to 2009, the phase with a large amount of precipitation corresponded to the phase of relatively low temperatures, and, conversely, the dry phase of the precipitation regime corresponded to relatively high temperatures (Obyazov, 2012).

The amazingly rich and diverse ecosystems of the Daurian steppe are concentrated in the Daursky State Nature Biosphere Reserve,
which facilitates the preservation of dozens of rare species and the study of many natural processes of the great steppe. Daursky State Nature Biosphere Reserve was created in 1987 to preserve and study the steppe, lacustrine-steppe and wetland complexes of Transbaikalia. The Daursky State Nature Biosphere Reserve is the only place in Russia where such unique species as the Relict Gull and the gazelle survive (Kirilyuk et al., 2009). The Daursky State Nature Biosphere Reserve forms part of a key bird area in Asia, being one of the most important resting places for millions of migratory waterfowl and semi-aquatic birds (Goroshko, 2000). The Daursky State Nature Biosphere Reserve, together with the Chinese and Mongolian reserves, forms an international Protected Area named «Dauria». Extreme dynamics and a low level of anthropogenic alteration to its natural systems are both important characteristics of the Daursky State Nature Biosphere Reserve (Kirilyuk et al., 2009).

The lakes Zun-Torey and Barun-Torey are the largest hydrological objects of the Daursky State Nature Biosphere Reserve zone. They are the remains of a large lake (sea) which once reached the size of 2400 km² (Shamsutdinov, 1983). Frish (1972) stated almost 50 years ago «In a few years, one can observe phenomena, the development of which in other places is known to exceed the human life span». Therefore, in the current context of the climate emergency, this area represents a testing ground for observing climatic changes and how various communities (terrestrial and aquatic) respond and adapt to these changes (Kuklin et al., 2013). A careful study of the nature of the Torey lakes (the standard for a variety of landscape transformations) may suggest how people can interact appropriately with the landscape in other regions (Frish, 1972). The pulsating water regime of the lakes determines the cyclicity of the terrestrial ecosystems (Kirilyuk et al., 2009, 2012) and periodic surges in the productivity of the aquatic ecosystems in the shallow soda lake (Afonina & Tashlykova, 2019, 2020; Bazarova et al., 2019).

The limnology of Torey soda lakes, despite their extensive area, has not been well studied. Even so, some authors have contributed to our knowledge of them (e.g. Ivanov, 1977; Alimov, 1991; Dzyuba & Kulagina, 2005; Tsyrenova et al., 2008; Namsaraev, 2009; Zakharyuk et al., 2010; Sklyarow et al., 2011). In this paper, we use long-term data from two lakes to document how changes in water level have affected physical and chemical variables in these water bodies. Our aim is to synthesise data on changes in the environmental parameters of the Torey lakes during a climate cycle, the drying and initial filling of basins. Studies of contemporary responses of lakes to climate shifts are of interest for planning and management (Olaka et al., 2010; Jeppesen et al., 2015). So this work is important and particularly important because the study area is becoming increasingly arid, and this trend is expected to continue under the current climate-change scenario (Obyazov, 1996).

### Material and Methods

#### Study site description

The Daursky State Nature Biosphere Reserve belongs to the Onon-Torey district of the dry Mongolian-Manchurian steppe, which is the northeastern part of the Central Asian physical-geographical region. Torey lakes are large in area and shallow; they constitute of one hydrological system and are connected by an Utochi channel (about 300 m long and 100 m wide) (Fig. 1).

A unique feature of the lakes Barun-Torey and Zun-Torey is the unstable hydrological regime. The reason for the water level change is the periodic filling and drying of lakes due to cyclic climatic changes in humidity and temperature. Over the past 200–220 years, the lakes repeatedly dried up and filled up with a frequency of about 30 years (Obyazov, 1996) (Fig. 2).

The water level regime of Lake Zun-Torey differs from that of Lake Barun-Torey, since its catchment area is smaller and it has no surface tributaries. When the water level is high, the hydrological regime in the lakes is similar, and they are connected by a channel. The runoff from Lake Barun-Torey to Lake Zun-Torey begins at the level of 596.1 m a.s.l. according to the Baltic system (BS). Once the water surface of the lakes is level up, the direction of the current in the channel changes under the influence of wind and other factors. When the water level falls, the shallower Lake Barun-Torey dries up quicker, and the water begins to move in the opposite direction through the connecting channel: from Lake Zun-Torey to Lake Barun-Torey. With a further fall in the water level the Utochi channel dries up and the lakes become isolated again (Obyazov, 1996).
In high water years with a water cut of 598 m a.s.l. BS, the surface area of Lake Barun-Torey reaches 550 km², the volume is 1.4 km³, the length is 40 km, the width is 20 km, the maximum depth is 4.3 m, and the average depth is 2.5 m. The catchment area is 25 700 km². There are approximately ten islands in the lake, and their number varies depending on the water level. The bottom of the lake is flat, and the deepest parts are in the centre of the lake. The shoreline is crenelated, with abundant headlands and bays. The shores are slightly swampy, often saline lands occupy (Kirilyuk et al., 2009, 2012; Kirilyuk & Tkachuk, 2012). According to Vakhnina et al. (2020), in 1999 the maximal lake area reached 557.8 km². Two rivers flow into the lake, of which the River Uldza flows up from the south and forms a vast estuary, while the River Imalka flows into the lake from the west. The main catchment area of the rivers is located in Mongolia. The outflow only occurs in high water years, while in dry years the riverbeds dry up. For this reason, and for the associated decrease in precipitation, the lake quickly becomes shallow, and its surface area shrinks, sometimes resulting in the lake drying up completely (Kirilyuk et al., 2009, 2012). By June 2009, Lake Barun-Torey had completely dried up while just before its floor contained only shallow large puddles, which appeared after rains (Kirilyuk et al., 2009). A trend was noted in 2013 that the lake was filling, in connection with the hiatus in the Pacific monsoon. In the summer of 2014, a flow was recorded in the River Uldza. But in the spring of 2015, the Lake Barun-Torey could not recover its water, and, as a result, the lake dried out again (Davydova, 2014). According to our observation, in the summers of 2014, 2016, 2018, 2020, the dry bed of Lake Barun-Torey was a steppe with some small water bodies (pools, puddles) (Fig. 3). In autumn 2020, we observed the discharge of a full-flowing River Uldza into the Lake Barun-Torey (Fig. 4).
Lake Zun-Torey is round, with a weakly indented shoreline and a single island, which turns into a peninsula when the water level lowers. When the water is high, the surface area of this lake is 285 km$^2$, the volume is 1.6 km$^3$, the length is 23 km, the width is 13 km, the maximum depth is 6.7 m, and the average depth is 4.5 m. The highest depth is in the northern part of the lake, and the lake bed drops steeply. The catchment area is 26,000 km$^2$ (Kirilyuk et al., 2009, 2012). In 1999 the maximal lake area was 301.6 km$^2$ (Vakhnina et al., 2020). When low water level, the only island connects to the mainland coast. There were wide mud bars, sandy and pebble beaches around the lake (Kirilyuk et al., 2012). In 2015, the lake area reached 254 km$^2$ (Vakhnina et al., 2020). Lake Zun-Torey completely dried out in autumn 2017. According to our observation, in 2018, the lake formed viscous impassable salt marshes with precipitation water, driven by winds to different lake shores. In 2020, the Lake Zun-Torey was a steppe where gazelle antelopes were grazing (Fig. 5).

Torey lakes are soda-chloride-sodium lakes with a high alkalinity. The lakes’ hydrochemistry is determined by a combination of an evaporative concentration of salt composition, the production of organic matter and by hydrogenic sedimentation (Zamana, 2009; Zamana & Borzenko, 2010).

Data sampling and analysis

We present a case study for the lakes Barun-Torey and Zun-Torey, which have been subject to major changes in water level. We have provided a brief overview of the characteristics of the lakes and add detailed information using unpublished data. We have been studying the lakes for a number of years, which we have provisionally divided into periods, including three periods for the Lake Zun-Torey (high (1999 and 2003), intermediate (2007, 2011) and low (2014 and 2016) lake levels) and two periods for Lake Barun-Torey (high water level (1999 and 2003)

Fig. 3. Lake Barun-Torey in various years. Designations: A – 1999 (Photo: G. Tsybekmitova), B – 2011 (Photo: G. Tsybekmitova), C – 2014 (Photo: P. Matafonov), D – pool 3 (puddle) in 2018 (Photo: B. Bazarova), E – pool 4 (Photo: B. Bazarova), G –2020 (pools 1 and 2 are visible in the distance) (Photo: Sh. Askarov).

Fig. 4. River Uldza in 2020 (Photo: E. Afonina).


In 2018 and 2020 (unpublished data), we used an Aquaread multiparametric sensor for water analysis (Great Britain) to determine the total of dissolved solids (TDS), pH, dissolved oxygen content, and water temperature. The abiotic parameters were measured at the surface of the pools. The water transparency was determined with a standard Secchi disk.

We used parametric and multivariate statistical analysis to classify the lakes according to the similarity of the major abiotic variables using XLSTAT Basic package (Addinsoft, USA). Variables included in the abiotic parameter matrix were normalised, with preliminary selective transformation of the initial values. Similarity in terms of sets of abiotic factors was estimated using normalised values of Euclidean distance based on hierarchical clustering and ordination (Dalgaard, 2008). We performed a principal correspondence analysis (PCA) for the water body groups and environmental variables using pooled sampling data.

**Results**

The first two components in the results of the PCA explained 67.67% of the variance in total (Table). The PCA performed with PCA divided the studied lakes and pools into four groups according water level and basic abiotic parameters (Fig. 6).

The first group of sites consists of lakes Barun-Torey and Zun-Torey in high water level, when depth ($r = -0.806$, $p < 0.001$) and transparency ($r = -0.827$, $p < 0.001$) were the highest. The second and third group include Lake Zun-Torey during intermediate and low water level respectively. In low lake water levels TDS ($r = -0.990$, $p < 0.001$) and pH ($r = 0.600$, $p < 0.001$) were the highest. In the fourth group, filling pools located on the Lake Barun-Torey dry bottom, water temperature ($r = 0.510$, $p < 0.001$) and dissolved oxygen contents ($r = 0.355$, $p < 0.001$) increased.

**Lake Zun-Torey**

In high water years (1999 and 2003), the maximal sampling depth was 6.5 m and 5.6 m respectively, the water transparency was 0.5 m. Abiotic parameters corresponded to TDS = 2.12 g × l$^{-1}$, pH = 9, $T = 20.4–23.4^\circ C$, $O_2 = 5.6$ g × l$^{-1}$.

In intermediate water level years (2007 and 2011), the maximal sampling depth was 2.3 m and 1.7 m respectively. The water transparency was 0.3 m. TDS increased to 5.1–8.2 g × l$^{-1}$, pH to 9.4–9.6, and water temperature to 23–24°C. Dissolved oxygen content changed from 4.8 g × l$^{-1}$ to 7.0 g × l$^{-1}$. In low water years (2014 and 2016), the maximal sampling depth was 1.5 m and 0.5 m respectively. The water transparency was 0.3 m. TDS increased to 5.1–8.2 g × l$^{-1}$, pH to 9.4–9.6, and water temperature to 23–24°C. Dissolved oxygen content increased to 7.8–10.3 g × l$^{-1}$ (Fig. 7). The water in the lake was an albescent colour (Fig. 5, Fig. 7).

**Lake Barun-Torey**

In 1999 and 2003 the maximal sampling depth was 4.5 m and 3 m respectively. The water transparency was 0.4 m and 0.3 m respectively. Abiotic parameters corresponded to TDS = 2.1 g × l$^{-1}$, pH = 9, $T = 20.5–24^\circ C$, $O_2 = 14.4$ g × l$^{-1}$. The water was an albescent colour (Fig. 3A, Fig. 8).

### Table. Results of component analysis of the summarised data

<table>
<thead>
<tr>
<th>Main components</th>
<th>Eigenvalue</th>
<th>Constrained inertia (%)</th>
<th>Cumulative (%)</th>
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<td>1</td>
<td>2.29</td>
<td>38.15</td>
<td>38.15</td>
</tr>
<tr>
<td>2</td>
<td>1.77</td>
<td>29.52</td>
<td>67.67</td>
</tr>
</tbody>
</table>
**Temporary pools on the Lake Barun-Torey bed**

Since 2014 at the northernmost point of Lake Barun-Torey dry bed (near the village of Kulusutai) we have observed some small water bodies (pools and/or puddles) with an area of 10–100 m² and a depth of about 0.1–0.5 m. In common the filling small temporary pools on the Lake Barun-Torey dry bed were characterised as alkaline (pH = 8.2–9.0), oligohaline (0.5–1.0 g × l⁻¹), well warmed up (26–30°C), and muddy (except pool 3) (TR = 0.1–0.3 m) (Fig. 8).

Temporary pool 1 and pool 2 are fed by a groundwater discharge through the fracture at the western shore of the lake. We observed aquatic plants (*Lemna* spp., *Potamogeton* spp., *Phragmites australis* (Cav.) Trin. ex Steud.) and filamentous algae in the pools. The bottom of the pools were primarily muddy and were homo-
Soda lakes with a high surface to volume ratio respond strongly to climatic changes (Olaka et al., 2010). In shallow lakes, the water level fluctuates naturally intra- and inter-annually depending mostly on the regional climatic conditions (Beklioglu et al., 2007). In this inland and fluctuating environment, acute changes in the physical and chemical environment occur throughout each hydrological cycle. Moreover, the duration of each cycle is different. Several aspects of the cycles have been studied, which alternate between wet and dry phases of inter-annual and seasonal environments (Jęppesen et al., 2015). In Southeastern Transbaikalia, in 1951–2009, the average annual temperature increased by 1.9°C. In different parts of the study area, the linear trend showed increases from 1.5°C up to 2.2°C during 59 years. It led to an increase of the period with positive temperatures from 165–167 days to 173–179 days. Long-term changes in precipitation are cyclic. In the study area, cycles are most vividly apparent within the time span of a century (Obyazov, 1994). From 1955 to 1963, there was a period with above-average precipitation, followed, up to 1982, by a below-average period. In 1983, a wet period started again, lasting until 1998. The beginning of a new dry period was in 1999. This period probably ended by 2012. Thus, up to two full cycles in precipitation occurred in the area over the past 60 years. The lake levels also directly depend on precipitation and correlate with river flows (Simonov et al., 2013). Changes in climate in the Daurian ecoregion, especially with regard to humidification cycles and continuing warming, caused alterations in habitats, and even habitat disappearance (Kirilyuk & Tkachuk, 2012). In the lakes Barun-Torey and Zun-Torey, the salt layer that has been formed after a period of drying is gradually becoming overgrown and turning into meadow and steppe. The lake depression is devoid of water, and instead of being an ecosystem inhabited by aquatic organisms and organisms that live near water, including thousands of fish, it has become part of the terrestrial ecosystem (Kirilyuk et al., 2012).

In Lake Zun-Torey, the salinity varied from oligohaline (2.1 g × l⁻¹ in 1999) to polyhaline (20.9 g × l⁻¹ in 2016), and the pH increased from 9.0 to 9.9 respectively during transitions from high to low water levels. The decrease in the heat capacity of the lake, caused by the increase in salinity, resulted in a rapid, strong heating of the water in the summer months (water temperature rose from 20.4–23.4°C to 21.4–26.4°C). The considerable area of Lake Zun-Torey at shallow depths promoted wind mixing of the water mass and its saturation with oxygen (dissolved oxygen increased from 5.6 g × l⁻¹ to 8.6 g × l⁻¹) (Afonina & Tashlykova, 2019).

About Lake Barun-Torey, unfortunately, we have no material to characterise the intermediate water level period. But according to Zamana & Borzenko (2010) in 2006 TDS was 9.7 g × l⁻¹ and the water type was characterised as mesohaline. The lake drain rather quickly after the river inflow ceases. Not far from springs, gushing out of the ground, there are small freshwater lakes and small ponds with bitter salt water. The exposed bottom of the lake (no longer underwater) is covered with a variety of rather sparse vegetation (Nekipelov, 1957; Simonov et al., 2013). An increase in atmospheric precipitation, due to the penetration of the Pacific monsoons into south-eastern Transbaikalia, along with the release of groundwater and the melting of permafrost, has led to the initial and partial filling of Lake Barun-Torey (Zamana & Vakhnina, 2014). The studied temporary pools located on the dry bed of Lake Barun-Torey show marked differences among one another, although they are all temporary, shallow and fresh systems (up to 1.0 g × l⁻¹). Temporary pool 1 and pool 2 are fed by groundwater discharge through the fracture at the western shore of the lake. The piezometric level of the studied water seepage is at a depth of 4 m and is above the earth surface (Borzenko et al., 2019, 2020).
The sensitivity of the water level to variations in climate depends on the bathymetry of lake basins, evaporation rates, and the interannual variability of water inflows (Wood & Talling, 1988; Talling, 2001). Periodic changes in the level of surface and ground water are noted in other endorheic shallow lakes in Russia (Vesnina et al., 2012; Vizer, 2016), China (Deng et al., 2001), Mongolia (Rasskazov et al., 1990; Sevastyanov, 2000), Europe (Comín et al., 1999; Legler et al., 2011), Africa (Schagerl, 2016), and North America (Stephens, 1990). In the dry period, the lakes completely dry up, while in the high-water period they are stable and filled with water. We would also like to note that among large alkaline water bodies vulnerable to climate change only one other lake, in Africa, has a 30-year hydrological cycle like the Torey lakes (Olaka et al., 2010; Schagerl, 2016).

Saline lakes in Central Asia provide numerous ecosystem services and have a high and often underestimated economic value (Zadereev et al., 2020). We believe that our work on the review of available literature and along with own data supports the sensitive and rare soda lakes. Soda aquatic ecosystems in arid regions are important habitats for aquatic, semiaquatic, and riparian wildlife. These habitats serve as a regional biodiversity source. Because of the variability of local, regional, and global climate, these lakes fluctuate in size and salinity and support a regional pool of species adapted to various levels of salinity (Kirilyuk et al., 2009; Kirilyuk & Tkachuk, 2012; Zadereev et al., 2020). Lake Barun-Torey and Lake Zun-Torey provide cultural ecosystem services. They have a symbolic meaning. Lakes are used for commercial fishery at the high water level. The Torey lakes are used as a recreational site for summer camping and fishing. A cattle grazing is forage base in the Torey lakes provision of ecosystem services. A reduced runoff, increased evaporation, and an increased climate variability form the most important challenges of climate change on the lakes.

Conclusions
In the regressive water phase of the climatic cycle (the drying and initial filling of basins), the water level falls in the lakes. This leads to changes in some physical parameters and hydrochemical compounds. In Lake Zun-Torey, we found an increase in such environmental characteristics as TDS (from 2.1 g × l−1 to 20.9 g × l−1), pH (from 9.0 to 9.9), dissolved oxygen concentration (from 5.6 g × l−1 to 8.6 g × l−1), and temperature (from 20.4–23.4°C to 21.4–26.4°C), and decrease in depth (from 6.5 m to 0.5 m) during transitions from high to low water levels. In a period of high water level, the water of Lake Barun-Torey is of the same constitution as in Lake Zun-Torey (oligohaline and alkaline). The temporary pools on the dry bed of Lake Barun-Torey are freshwater, arising from groundwater and precipitation.

Acknowledgements
We would like to thank our colleagues at the Laboratory of Aquatic Ecosystems (Institute of Natural Resources, Ecology and Cryology of SB RAS, Russia) for their assistance in collecting samples. We are grateful to the staff of the Daursky State Nature Biosphere Reserve (V.E. Kirilyuk, O.K. Kirilyuk, O.A. Goroshko, A.A. Cherepysyn, A.P. Borodin) for organising field research. We thank P.V. Matafonov, G.Ts. Tsymbekmitova, B.B. Bazarova, Sh.A. Askarov for some photos provided. This study was supported by the Program for Basic Research of the Siberian Branch of the Russian Academy of Sciences, project no. FUF-2121-0006. Authors are grateful to the anonymous reviewers and editors for having substantially improved this manuscript with their comments and advice.

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ТОРЕЙСКИЕ ОЗЕРА (ДАУРСКИЙ ГОСУДАРСТВЕННЫЙ ПРИРОДНЫЙ БИОСФЕРНЫЙ ЗАПОВЕДНИК, РОССИЯ): МНОГОЛЕТНИЕ ИЗМЕНЕНИЯ ПАРАМЕТРОВ СРЕДЫ

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Ключевые слова: абиятические показатели среды, Барун-Торей, гидрологический цикл, Даурский государственный природный биосферный заповедник, Зун-Торей, уровень воды